Holocene fire and occupation in Amazonia: records from two lake districts

Mark B. Bush1,*, Miles R. Silman2, Mauro B. de Toledo1,4, Claudia Listopad1, William D. Gosling1, Christopher Williams1, Paulo E. de Oliveira3 and Carolyn Krisel2

1Department of Biological Sciences, Florida Institute of Technology, Melbourne, FL 32901, USA
2Department of Biology, Wake Forest University, Winston Salem, NC 27106, USA
3Department of Biological Sciences, Universidade do Guarulhos, Sao Paulo 07023-070, Brazil
4Marine Geology and Geophysics Department, Universidade Federal Fluminense, Niterói RJ 24.210-340, Brazil

While large-scale pre-Columbian human occupation and ecological disturbance have been demonstrated close to major Amazonian waterways, less is known of sites in terra firme settings. Palaeoecological analyses of two lake districts in central and western Amazonia reveal long histories of occupation and land use. At both locations, human activity was centred on one of the lakes, while the others were either lightly used or unused. These analyses indicate that the scale of human impacts in these terra firme settings is localized and probably strongly influenced by the presence of a permanent open-water body. Evidence is found of forest clearance and cultivation of maize and manioc. These data are directly relevant to the resilience of Amazonian conservation, as they do not support the contention that all of Amazonia is a ‘built landscape’ and therefore a product of past human land use.

Keywords: agriculture; charcoal; fossil pollen; Peru; Brazil; pre-Columbian

1. INTRODUCTION

The role of pre-Columbian human activity in shaping Amazonian ecosystems has received considerable attention in the light of occupational records from the Xingu River basin, Marajos Island, and near the modern cities of Santarém and Manaus (Roosevelt 1991; Roosevelt et al. 1991; Heckenberger et al. 1999). Advocates of widespread human influence in Amazonia point to what appear to be urban centres (Roosevelt 1991; Roosevelt et al. 1991; Heckenberger et al. 1999, 2003), large earthworks in Bolivian savannahs bordering Amazonia (Mann 2000; Erickson 2001), and the widespread occurrence of soils enriched with carbon (terra preta) within Amazonia (Glaser et al. 2001; Lima et al. 2002). If Amazonia was a managed landscape prior to European contact, the biodiversity of the region has withstood substantial fragmentation, hunting and human interaction, and such manifest resilience would influence conservation decision making. While Erickson (2001) has described the Bolivian savannah landscape as having been manufactured by humans, other authors are generally more circumspect when dealing with the densely forested regions.

Major uncertainties exist in the size of the pre-Columbian human population of Amazonia, with estimates ranging from 1 to 11 million inhabitants. Similarly, the large areal extents of some archaeological settlements, e.g. 5 km² of terra preta underlying the modern city of Santarém, have been variously interpreted as representing continuous occupation by a large population (Roosevelt 1987) or repeated resettlement and abandonment by smaller groups (Meggers 1995). Heckenberger et al. (2003) have argued for dense settlement and societies with strong social hierarchies with the capability to transform landscapes. The alternate view is that human endeavours in Amazonia were constrained by poor soils and that occupation was sparse, lacking the societal complexity or monuments of Andean or Central American systems (Meggers 1954, 2003a,b). Denevan (1996) suggested that occupation was centred on sandy bluffs overlooking navigable channels and floodplains. This interpretation of existing data suggests local alteration of landscapes extending 5–10 km around settlements, but with vast interfluvial areas where humans had little impact on the biota (Denevan 1996).

From an ecological standpoint, humans are likely to deplete animal populations and transform landscapes through burning. While it is very difficult to document the past population size of people, let alone animals, evidence of burning is well preserved in palaeoecological records. Indeed, we suggest that fire is the characterizing disturbance of human activity in much of western Amazonia and we will take its presence to be an indicator of human activity.

(a) Fire in Amazonian landscapes

Human settlers modified the landscape through fire. Some parts of Amazonia are more susceptible to fire...
than others (Nepstad et al. 2004) and considerable uncertainty exists over the ‘natural’ occurrence of Amazonian forest fires. Some researchers have found evidence of natural fire (e.g. Sanford et al. 1985; Saldarriaga & West 1986; Piperno & Becker 1996), whereas other data point to fire as very rare in the absence of humans (e.g. Hammond & ter Steege 1998; Turcq et al. 1998; Behling & Da Costa 2000; Behling 2001; Behling et al. 2001). In a 50 000-year palaeoecological record from the Hill of Six Lakes in northwestern Amazonia, only one sedimentary layer at 5600 calendar years before the present (yr BP) was found to contain significant amounts of charcoal, suggesting the rarity of fire in that landscape (Bush et al. 2004).

Piperno & Becker (1986) recorded charcoal in soils around Manaus with a concentration of ages between ca 900 and 1260 yr BP, and attributed these fires to natural climatic change rather than agriculture. From a nearby site, Santos et al. (2000) also recovered charcoal from soils with peaks between 1000 and 1300 yr BP. The period ca 1100–1300 yr BP was a phase of intensified El Niño activity (Moy et al. 2002), making it plausible that extreme drought caused fire at this time. However, some caution is needed in interpreting these data. Rather than slash and burn agriculture, less radical land management involving fire may have prevailed, e.g. that practised for Brazil nut harvesting in Peru (Phillips 1993) and Caryocar edule fruits by indigenous peoples in the Xingu (P. E. de Oliveira 2005, personal observation).

Nepstad et al. (2004) document that the forests of central Amazonia are only susceptible to natural fire during the late stages of the dry season. Under modern conditions in strongly seasonal forests, there is about a 10-day window each year when natural fire is possible. Near Manaus, the probability of fire decreases exponentially with distance from sites of human occupation (Laurance & Williamson 2001). In Peruvian Amazonia, indigenous people find it difficult to burn the forest at any time.

Evidence of the rarity of fire in western Amazonia comes from historical and palaeoecological observations. Cocha Cashu in Manu National Park lies less than 200 km from the western Lake District and has been continuously monitored for the past 40 years. In that time, no fire has been documented within the approximately 50 000 km² reserve. The only charcoal recovered from the Cocha Cashu lake sediments were fragments derived from wood burned since 1968 by the lake-shore field station (M. R. Silman, unpublished core data). Sediment cores raised from three oxbow lakes along the Madre de Dios, all of which span the dry season which ranges from two to four months in duration. The forests in the region are upper-Amazonian rainforest and have been extensively documented, e.g. Pitman et al. (2001) and are rich in Calycophyllum, Calophyllum, Cecropia, Cedrela, Ceiba, Diperyx, Ficus (35 species) and Poulsenia (Foster 1990). The palms Iriartea deltoidea and Astrocaryum murumuru are the commonest stems in the forest (Pitman et al. 2001). Of note is the local abundance of Lecythidaceae (particularly Bertholletia excelsa and Couratari spp.) in the forests. Temperatures are relatively constant with a mean of 25°C, but the lake district lies across a region of sharply changing precipitation ranging from approximately 2000 mm yr⁻¹ in the south to approximately 1700 mm yr⁻¹ in the north. Similarly, there is some variability in the dry season which ranges from two to four months in duration. A modern road passes closest to Lake Gentry, and on a knoll in lake records in the absolute concentration of charcoal in lake sediment. However, as lake size and taphonomic processes influence rates of charcoal deposition (Whitlock et al. 1997; Whitlock & Larsen 2001), comparisons of concentration to infer intensity of land use are most robust within a lake record, rather than between lake records.

Although landscape alteration is clearly going to have occurred around known archaeological sites, the extent to which those sites are representative of other portions of Amazonia is unknown. Consequently, the spatial and temporal scale of landscape alteration in previously undocumented settings may provide an alternative image of Amazonian occupation. We present data from two Amazonian lake districts that document human occupation, and use the charcoal and pollen records from those sites to gauge the spatial and temporal extent of landscape alteration at those locations. In particular, we ask ‘did human occupation result in widespread landscape modification at these two widely separated Amazonian sites?’ or, as hypothesized by Denevan (1996), is human disturbance centred on particular features of the landscape?

2. MATERIAL AND METHODS

(a) Study areas

A paucity of long Holocene palaeoecological records exists in Amazonia, but here we report on two suites of cores raised to investigate climate change in ecologically sensitive areas of Amazonia (Listopad 2001; M. B. de Toledo 2004, unpublished Ph.D. thesis). Sites were selected to provide lakes that were not directly influenced by rivers, lay within terra firme forest and where replication could be achieved within a small geographical area. A western lake district (hereafter western lakes) near Puerto Maldonado, Peru, that comprised lakes Gentry, Vargas, Parker and Werth, and an eastern lake district (hereafter eastern lakes) near Prainha, Brazil, that comprised lakes Geral, Santa Maria and Saracuri, were selected for study. The western lakes lie within a lake district containing, approximately 50 lakes and swampy depressions (approx. 12.5° S, 69.0° W; figure 1a). The lake district lies at approximately 230 m elevation and all of the lakes are relatively small, approximately from 0.3 to 2 km in diameter and are approximately 1–3 m deep. The origin of the lakes is unknown, but they lie at the highest points on the rolling landscape, outside fluvial influence, and are underlain by sands and clays.

The lakes lack inflowing streams and their permanence is indicated by the presence of fishes. The lakes appear to be mesotrophic and are fringed by beds of Sagittaria, Pontederia and Eichornia. The forests in the region are upper-Amazonian rainforest and have been extensively documented, e.g. Pitman et al. (2001) and are rich in Calycophyllum, Calophyllum, Cecropia, Cedrela, Ceiba, Diperyx, Ficus (35 species) and Poulsenia (Foster 1990). The palms Iriartea deltoidea and Astrocaryum murumuru are the commonest stems in the forest (Pitman et al. 2001). Of note is the local abundance of Lecythidaceae (particularly Bertholletia excelsa and Couratari spp.) in the forests. Temperatures are relatively constant with a mean of 25°C, but the lake district lies across a region of sharply changing precipitation ranging from approximately 2000 mm yr⁻¹ in the south to approximately 1700 mm yr⁻¹ in the north. Similarly, there is some variability in the dry season which ranges from two to four months in duration. A modern road passes closest to Lake Gentry, and on a knoll...
In addition to the three lakes (Vargas, Parker and Gentry) formed between 8000 and 6000 yr BP, this was not a uniformly wet period. Consistent with other regional data (Bush et al. in press), a mid-Holocene dry event influenced the region, but was interrupted by wet episodes, presumably when the lakes filled (Paduano et al. 2003). The lakes show different levels of sensitivity to drought, with Vargas most sensitive, then Parker, and Gentry the least sensitive. However, even Gentry was susceptible to drying and prior to ca 4800 yr BP exhibited at least two brief phases when pollen was oxidized from the sediments. Another period of apparently slow sedimentation between ca 1270 and 2800 yr BP may point to intermittent accumulation, although high pollen concentrations suggest that this may simply have been a time of little sediment input; otherwise the Gentry record appears to have a continuous history of pollen accrual. The last lake to form was Werth, which started to accumulate sediment ca 3400 yr BP, and deepened ca 900 yr BP.

Once the lakes had formed fully, they provide pollen records with greater than 70% arboreal pollen throughout their history (figure 2). Only Gentry showed evidence of agriculture, with Zea pollen found between 3700 and 500 yr BP and two grains of Manihot pollen at ca 2400 yr BP. Charcoal was most abundant in Gentry, ranging between 1 and 34 mm$^2$ cm$^{-3}$ throughout much of the record. Notably, charcoal was absent during the second of the oxidizing phases at ca 5000 yr BP.

Lake Parker also contained charcoal through much of its history, including the time immediately prior to the formation of Gentry. Charcoal concentrations in Parker were generally about half those of Gentry, between approximately 0.5 and 17 mm$^2$ cm$^{-3}$ (figure 2). Lake Vargas contains no charcoal for most of its history, except a strong spike of charcoal with approximately 17 mm$^2$ cm$^{-3}$ at ca 1000 yr BP. The Lake Werth record contains no charcoal except for three tiny fragments at ca 800 yr BP.

(c) Eastern lakes

The eastern lakes, Saracuri, Santa Maria and Geral all formed between 8200 and 8400 yr BP. The previously reported history for Geral (2000; Piperno & Pearsall 1998) is substantiated and supplemented by the new analyses.

The small basin of Santa Maria provides a sensitive proxy for lake level change. High modern aquatic pollen abundances are due to a 40 m wide floating mat of marshland species. Poaceae pollen rises and falls synchronously with the aquatics, indicating that at this site Poaceae pollen are probably derived from the marsh fringe. Notably, Poaceae percentages range as high as 85% at Santa Maria, whereas they seldom rise above 12% at the other two lakes. Saracuri provides a remarkably stable pollen record throughout its history (figure 3) and

IAS 1000–3000
Mauritia and Mauritiella were common on shorelines and in swamps. Mean annual temperature is 27°C and rainfall is approximately 2200 mm per annum over an eight and nine month wet season (Instituto Brasiliero de Geografia e Estatistica 1990). Though this climate could support tropical semideciduous forest, the local vegetation is a savannah–forest mosaic. The origin of the savannah is unknown and could be the result of edaphic factors or a long history of human landscape alteration (Prance & Schubart 1977).

Saracuri is completely enclosed by forest, with just some recent deforestation on the northern shoreline. Parts of the Santa Maria catchment are recently deforested, but a patch of tall dry forest lies within its catchment. Geral is an elongated lake running approximately east–west. Along its northern shore is forest, with savannah along its southern shore. Today, Geral is extensively used for leisure and fishing, receiving evidence of human use. The fossil pollen and phytolith record appears to have a continuous history of Geral (2000; Piperno & Pearsall 1998; Bush et al. 2000).

(b) Field and laboratory techniques

Between 1990 and 2001, cores were raised from the centre of each lake using a Colinvaux-Vohnout coring rig from a raft of rubber boats. The core from Geral was initially analysed at the Smithsonian Tropical Research Institute, Panama, and later, each lake using a Colinvaux-Vohnout coring rig from a raft of...
contrasts with the post-4000 yr BP portion of the Geral record, lacking high proportions of Poaceae, high charcoal content and the pollen from known cultivars. The divergence of the Poaceae curve from Geral with that of Saracuri is consistent with the onset of agriculture and probably reflects human activity around Geral.

Geral is the only lake of this trio in which Zea mays pollen was found. Pollen and phytoliths of Zea are found consistently after ca 4030 yr BP and prior to ca 850 yr BP. A phase of more active forest burning is evident in both the local charcoal reported here (figure 3) and a doubling of the concentration of charred Heliconia phytoliths at this level (Piperno & Pearsall 1998). Clearly, while cultivation of maize may have taken place on the shoreline and in seasonally exposed shallows (Roosevelt 1980; Matheny & Gurr 1983), there was also clearance of adjacent forest. Santa Maria and Saracuri both contained charcoal, including samples predating the first evidence of charcoal at Geral but at lower concentrations than in Geral. Within the period of known crop cultivation, charcoal concentrations at both Santa Maria and Saracuri fall from their mid-Holocene highs, whereas those of Geral show a much lesser decline. Indeed, the early Holocene peaks of concentration are slightly higher at Santa Maria (as befits a small basin) than at Geral, but fall to about one-fifth those of Geral post-4000 yr BP, suggesting much lower fire frequency than at Geral.

(d) The charcoal record

A regional drought influenced western and southern Amazonia between ca 6200 and 3400 yr BP (Servant et al. 1981; Mayle et al. 2000; Listopad 2001; Mayle & Beerling 2004), though its occurrence in central Amazonia is much less clearly defined (Bush et al. 2000; Behling et al. 2001; Listopad 2001; M. B. de Toledo 2004, unpublished Ph.D. thesis). During this period, fire frequency in the western lakes does not appear to be markedly different to other periods of the Holocene even though drought causes lake level to fall, exposing the mud–water interface. While organic material fails to accumulate and pollen is oxidized, charcoal continues to accumulate as it oxidizes very slowly. In none of the lakes is there evidence of a charcoal spike associated with a time of low pollen concentrations. Fires in these forests probably accompanied human activities and dry episodes caused site abandonment. At Gentry, the few samples that contained no pollen also contained no charcoal. Similarly, the record from Vargas overlaps with that of Parker between ca 7500 and 7000 yr BP and shows no peak of charcoal associated with the highest proportion of Poaceae. We infer that fire was rare or absent from this setting in the absence of human disturbance.

Vargas was the first lake to form and appears to have been briefly occupied and then abandoned. The abandonment is approximately coincidental with the colonization of Lake Parker and the first charcoal evident in that record. As Gentry filled, it was quickly occupied and the fire intensity and frequency became greater than that of any of the other sites.

In the eastern lakes, all three lakes contain some charcoal, but the lake that yielded crop pollen had the most consistent occurrence of relatively high concentrations of charcoal. Although Geral shows a more consistent fire history than the other two lakes, even the forested sites of Santa Maria and Saracuri were prone to fire, especially in the early Holocene. The highest charcoal peaks in those records are between 6900 and 7500 yr BP. The Geral record is not as well dated, but a peak of charcoal 6700 and 7200 yr BP is seen to overlap with that of Parker between ca 7500 and 7000 yr BP. The charcoal record does not overlap with that fire peak at other sites.

At Saracuri and Santa Maria, diminishing charcoal inputs indicate that fire becomes relatively rare ca 6800 yr BP. Between 6200 and 4500 yr BP, an oscillation in the aquatic and Poaceae pollen signatures, and a spike of charcoal at Santa Maria, suggest...
the possibility of a somewhat drier or more fire-prone system, although the lake does not dry out. At Geral, increased charcoal and Cecropia abundance characterize this period, suggesting increased disturbance and gap formation in the local forests, probably attributable to human activity. While this evidence clearly indicates human activity, the first occurrence of landscape alteration probably dates to at least 7700 yr BP, i.e. the first occurrence of charcoal in the record. One possibility is that land use intensified around 6000 years ago, coinciding with the increase in Cecropia pollen and the nearly...
contemporaneous occurrence of burned phytoliths and Poaceae pollen. However, it can also be argued that prior to the formation of an open body of water at the coring site, the deposition of microfossils may have reflected so local a source that land use beyond the immediate vicinity of the sample site cannot be detected. Hence, we cannot specify a time for the onset of disturbance at Geral.

The scale of impact
Prior to what may have been a single local fire event at Vargas, the scale of occupation appears to be consistent with modern anthropological observations that indigenous people manipulate relatively small territories. The actively managed area is much smaller than the hunting range. Studies of the Secoya–Siona in Ecuador revealed an occasional hunting range as large as 2500 km², with 590 km² being heavily used (Vickers 1988). The Secoya–Siona hunters live in the black-water systems of the Cuyabeno River, and hunting ranges in the fertile white-water systems of Peruvian Amazonia might be smaller. The Tsimane of northern Bolivia concentrate their farming and hunting activity within 3 km of their settlement (Apaza et al. 2002). Carneiro (1970) has suggested that a radius of approximately 5 km could support a sedentary population of 500 through swidden agriculture indefinitely. We have no way to estimate...
population size around Geral or Gentry. If we assume that these lakes were the centre of occupation, it is evident that the strength of human influence on the landscape decreased markedly with increasing distance from them.

After several thousand years of apparent use, the first palynologically identifiable crop is found in the sediments of Gentry. Only Gentry provided direct evidence of cultivation with pollen of *Zea* being found regularly between 3700 and 500 yr BP and manioc at ca 2400 yr BP. There is no clear spike in charcoal associated with the onset of maize cultivation; therefore, the field systems were already cleared or the cultivation took place on exposed mud when lake levels were low. The onset of maize cultivation coincides with a period of very slow sedimentation at Gentry, consistent with intermittent sediment accumulation and lowered lake levels, offering the opportunity to cultivate exposed wetlands. Detection of maize pollen at this time may not be a coincidence, as such large pollen grains (greater than 90 μm) are poorly dispersed, and cultivation closer to the sampling point improves the probability of recovering the fossil grain. Taking the charcoal and crop data together, we infer that humans have occupied this region of Amazonia for more than 8000 years, and through the use of fire they have altered the landscape at least at a local level. From our data, we cannot determine whether the occupation was seasonal or permanent, though the abundance of pottery shards on the ridge beside Gentry suggests a degree of permanence.

The sequential occupation of Vargas (and its abandonment), Parker and Gentry suggests that inhabitants were continually taking advantage of better home sites and may have abandoned these sites during the peak drought event at ca 5000 yr BP. If we are correct that Gentry was the centre of the occupation, we can compare the relative impacts to Parker, Vargas and Werth in terms of distance from the primary settlement.

The distance between Gentry and Parker is 9.5 km and between Gentry and Vargas is 17.5 km. Parker appears to have been used fairly extensively, but perhaps less intensively than Gentry, as no evidence was found of crop cultivation. Only in the last millennium were Gentry, Parker and Vargas simultaneously occupied. This expansion is coincident with the major Xingu landscape transformation (Heckenberger et al. 2003), but we note that it did not extend to the more remote setting of Lake Werth, 50 km from Gentry, which was never impacted by human occupation.

A similar image emerges from the eastern sites. If it is assumed that the long history of agriculture at Geral indicates that this lake formed the centre of local occupation, the distance to the adjacent lakes provides a scale of local impact. Saracuri at 5 km distance and Santa Maria at 6.7 km distance were both lightly used. No evidence exists to suggest that sites more than 5 km from the centre of occupation were heavily modified, though undoubtedly they would have been the site of hunting activity. Although the eastern sites lie within 150 km of the major pre-Columbian settlements around Santarem, that pattern of large landscape transformation is not seen in these records.

Our data are entirely consistent with the observation by Smith (1980) that the average size of terra firme terra preta deposits was approximately 1.4 ha, whereas riverside sites averaged 21.2 ha. This observation supports a history of patchy disturbance, with riverside settings not being typical of settings remote from rivers. Smith (1980) specifically noted that sites near the confluence of the Rio Negro and the Solimoes rivers and on the Xingu were particularly large. Consequently, we conclude that while disturbance was probably profound at a landscape level around major settlements, such localities should not be used to generalize about the state of Amazonia as a whole.

4. CONCLUSIONS
We offer the first landscape-scale view of disturbance at two sites in Amazonia and see long histories of disturbance at what appears to be a central location, with relatively small impacts on nearby systems and no apparent impacts on a system 50 km away. While these data do not support the contention of widespread landscape alteration, we note a considerable increase in human alteration of the landscape of the western lakes in the millennium prior to abandonment. Clearly, palaeoecological evidence exists for a population collapse at many sites across Amazonia and in isthmian regions of Central America (e.g. Bush et al. 1989; Piperno 1990; Bush & Colinvaux 1994).

In our experience of studying more than 30 lake records from Amazonia, a general rule of thumb has emerged: if the lake has modern use, has a road to it or has a settlement beside it, there is a very good probability of finding a long record (thousands of years) of disturbance. If, on the other hand, the lake is remote from modern society, it probably has a history lacking the characteristic signature of human activity. Denevan’s (1996) bluff model of occupation often coincides with the sites of modern occupation and hence our observations align well with his model. Although exceptions can always be found, modern use of an area is a powerful predictor of past human activity.

We observe that the extent of deforestation around the lakes that we discuss in this paper, and also the minimally seasonal locations of Kumpak (Liu & Colinvaux 1988) and Ayauchi (Bush et al. 1989) in Ecuador, was probably considerably less than around lakes in Central America with comparable histories of occupation, e.g. La Yeguada (Piperno et al. 1991) and Lake Wodehouse (Bush & Colinvaux 1994). It may be that the sites with very strong seasonality, which includes the Xingu, were more extensively impacted than less seasonal settings. We suggest that extrapolating from known areas of intense occupation to infer that most of Amazonia was a parkland in 1492 AD (458 yr BP) would be unwise. Rather we suggest that the highly heterogeneous history of land use, exploitation levels and population densities in Amazonia should be recognized, as these probably varied widely across this vast region in response to local soils, climate and potential protein sources. We now know that tropical forests have been subject to considerable climate change throughout the Quaternary. Forests of the Holocene are different in relative species abundance and productivity than they were during much of the Pleistocene. And even within the Holocene, natural
disturbance cycles such as droughts and disease may have played an important role in reshaping communities.

Hence, pre-Columbian Amazonian landscapes that differ from those of today could have been due to both direct and indirect human influence, natural causes, or most likely, some combination of all of these. Direct human actions include land conversions such as felling trees and setting fires. The temporal and spatial scales of these disturbances need to be resolved across the range of Amazonian landscape types. Indirect actions include hunting, introduction of exotic species that then modify a landscape or a disease that crosses from one host to another. Megafauna have large and clear landscape-level effects on forest structure and composition in systems where they still exist. What was the effect of their extinction on Amazonian forests? If, as has been suggested, the loss of megafauna was attributable to humans, then the case for a human modification of Amazonian landscapes is strengthened, although not on the time- or spatial-scale of the current archaeological debate.

Natural disturbance in the Holocene of Amazonia was probably also profound. While there is increasing documentation of millennial scale droughts (see §3d), contemporary studies indicate the potential importance of unexplained changes in animal abundance. In the 1980s, white-lipped peccaries disappeared, apparently stochastically, from an area of more than 70,000 km² for 12 years. The absence of this major seed predator resulted in a surge in recruitment of the dominant rainforest tree species (Silman et al. 2003, Wyatt & Silman 2004), an imprint that will last for decades, perhaps centuries, on this region.

Two potentially crucial differences exist between pre-Columbian human and natural disturbance in many areas of Amazonia. The first is the scale of disturbance. Modern studies of fragmentation show that duration and areal extent are critically important to post-abandonment succession (Laurance et al. 2002). Were the landscapes created by pre-Columbian occupants of Amazonia on the scale of a large tree-fall gap, or orders of magnitude larger? How long did disturbance persist on the same location? The second is the introduction of fire. Repeated burning transforms rainforest communities, whereas a single fire has lesser effects (Cochrane & Schulze 1998; Haugaasen et al. 2003). Determining whether the same patch of forest was burned repeatedly or whether there were long periods of recovery between events is an important question that needs to be resolved before making predictions of human impacts in any locale.

The conservation message that emerges from this and other palaeoecological studies is that many forests may be only one or two tree generations away from a period of land use and that in such areas populations are unlikely to be equilibrial (e.g. Foster et al. 2002). However, the type of mixed land use of the past was utterly different to the deforestation that accompanies modern ranching and soyabean farming. Away from the largest settlements, it is probable that gaps created by pre-Columbian farmers were closer to the scale of natural blowdowns, i.e. 1–10 ha, rather than the thousands of hectares created by agro-industry. Consequently, seed sources remained locally available and biodiversity was not reduced. Modern biodiversity and forest composition are not a product of land management, but have persisted despite it.

We are grateful to Paul Colinvaux for his role in directing the coring of Lake Gentry and to Chengyu Weng, Emilio Ancaya and Matthew Scriper for coring lakes Parker and Vargas. Carol Mitchell and Rolando Soto are thanked for logistical support in Peru. Hermann Behling and an anonymous reviewer provided thoughtful comments that strengthened this paper. This work was funded by NSF grants DEB 9732951, 0237573 (M.B.B) & 0237684 (M.R.S) and CNPq (doctoral scholarship 200065/99-8) (M.B.T).

REFERENCES


Phil. Trans. R. Soc. B (2007)


